

ห้องปฏิบัติการวิจัยการติดตามและการจัดการทางอุทกวิทยาเกษตรด้วยระบบอัจฉริยะ

Research Laboratory for INtelligent system

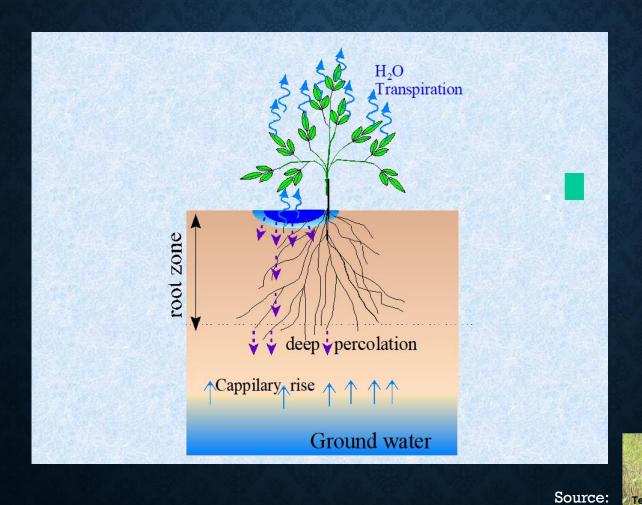
in Agro-hydrological Monitoring and Management (INAMM)

Department of Irrigation Engineering, Faculty of Engineering at Kamphaeng Saen, Kasetsart University

Soil-Water-Plant Relationship

Ekasit Kositsakulchai

Training on Climate Smart Irrigation Technology, 22 May 2018



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Irrigation & Fertilization Consultant

Soil Water

Soil properties

Water retention in soil

Infiltration

Soil Properties

Definition, soil components, soil profile and soil horizon, soil texture, soil structure, soil density and soil specific gravity, pore space



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Soil

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For other uses, see Soil (disambiguation).

Soil is a mixture of organic matter, minerals, gases, liquids, and organisms that together support life. The Earth's body of soil is the pedosphere, which has four important functions: it is a medium for plant growth; it is a means of water storage, supply and purification; it is a modifier of Earth's atmosphere; it is a habitat for organisms; all of which, in turn, modify the soil.





soil1

/soil/ •

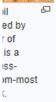
noun

the upper layer of earth in which plants grow, a black or dark brown material typically consisting of a mixture of organic remains, clay, and rock particles.

"blueberries need very acid soil"

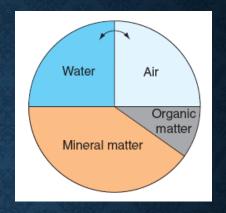
synonyms: earth, loam, dirt, clay, gumbo; ground "acid soil"

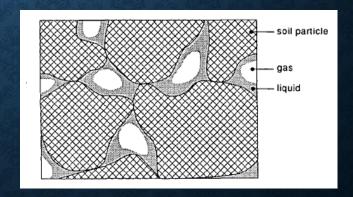
 the territory of a particular nation.
 "the stationing of U.S. troops on Japanese soil" synonyms: territory, land, domain, dominion, region, country
 "Canadian soil"



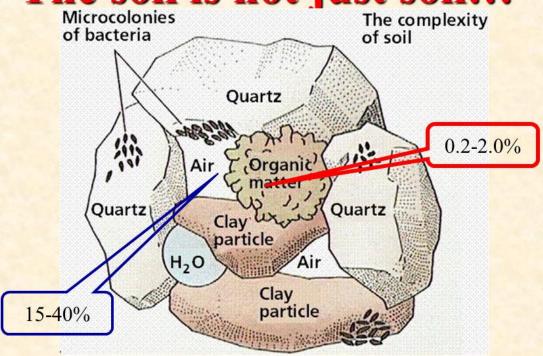
Soil Components

- Soil as a three-phase system
- Solid
 - Inorganic materials
 - Organic materials
- Liquid
 - Water
 - Solution
- Gas
 - Soil pores





The soil is not just soil... Microcolonies The complexity

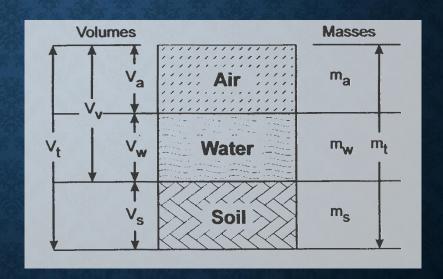


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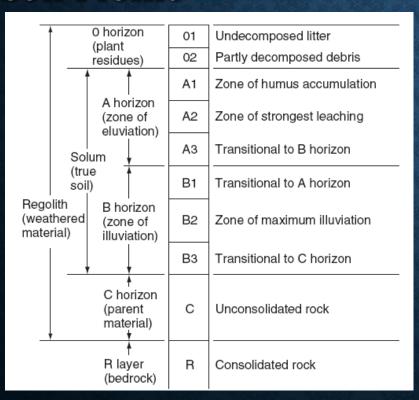
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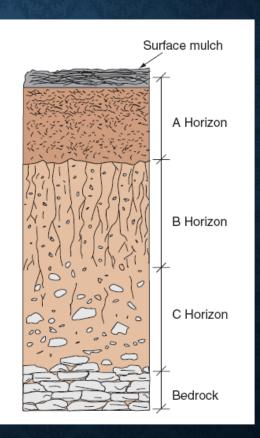
Soil Component

- Vt:total volume
 - Vv : void volume
 - Va: air volume
 - Vw: water volume
 - Vs : soil particle volume
- mt : total mass
 - ma: air mass
 - mw: water mass
 - ms : soil particle mass



Soil Profile







Comparison of two soil profile from different climatic region



Semi arid zone



Rainy region



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Soil Formation

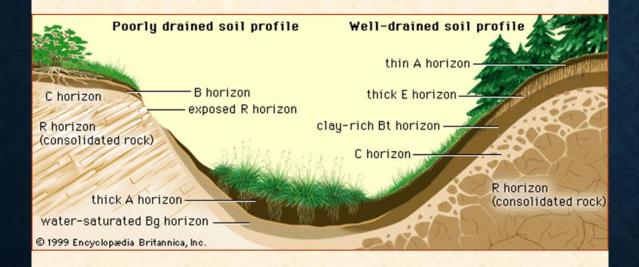


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Soil Formation



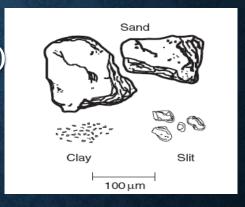
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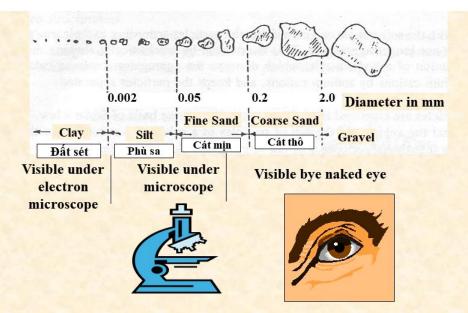
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Soil Texture

- A physical property of soil
- Defined from fraction of mass of soil particle (smaller 2 mm)
 - Sand: 0.05 2.00 mm
 - Silt: 0.002 0.05 mm
 - Clay: < 0.002 mm



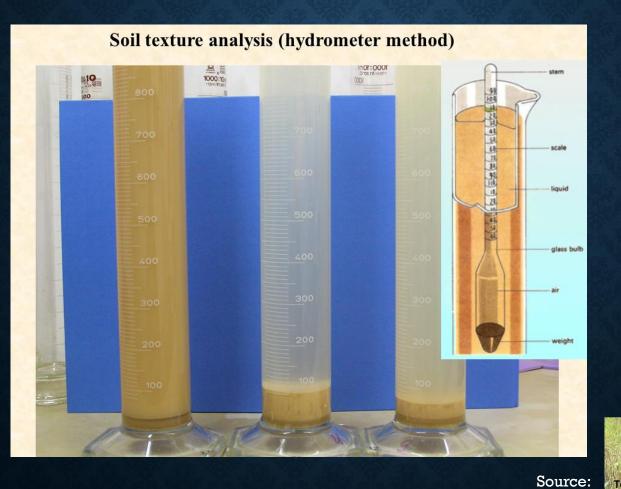
The size fractions in the soil



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Soil particle classification system

- USDA
- USPRA
- BSI, MIT
- ISSS
- DIN

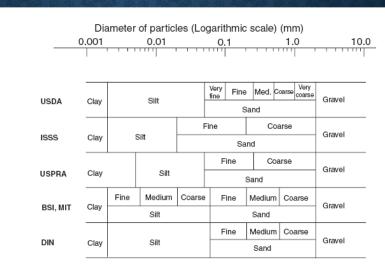
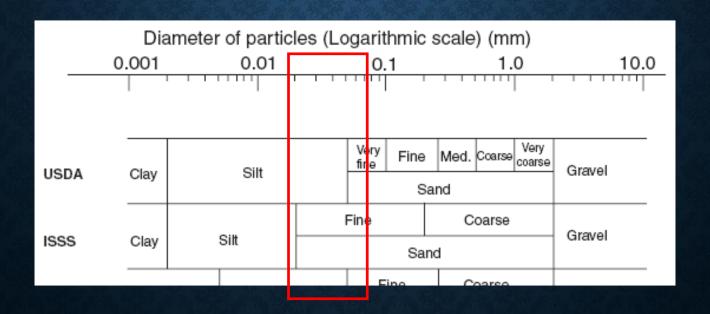


Fig. 3.1. Several conventional schemes for the classification of soil fractions according to particle diameter ranges; U.S. Department of Agriculture (USDA); International Soil Science Society (ISSS); U.S. Public Roads Administration (USPRA); British Standards Institute (BSI); Massachusetts Institute of Technology (MIT); German Standards (DIN).

Soil texture classification

- USDA U.S. Department of Agriculture
- ISSS International Soil Science Society



Soil texture classification

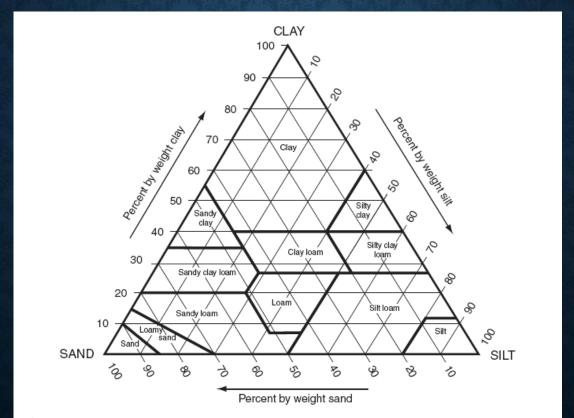
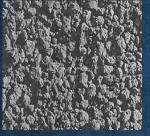


Fig. 3.3. Textural triangle, showing the percentages of clay (below 0.002 mm), silt (0.002–0.05 mm), and sand (0.05–2.0 mm) in the conventional soil textural classes.

Soil Structure





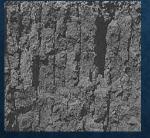


Granular



Blocky









Prismatic

Columnar

Soil Structure

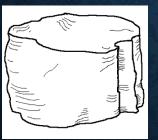


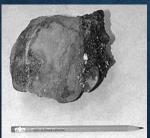




Platy

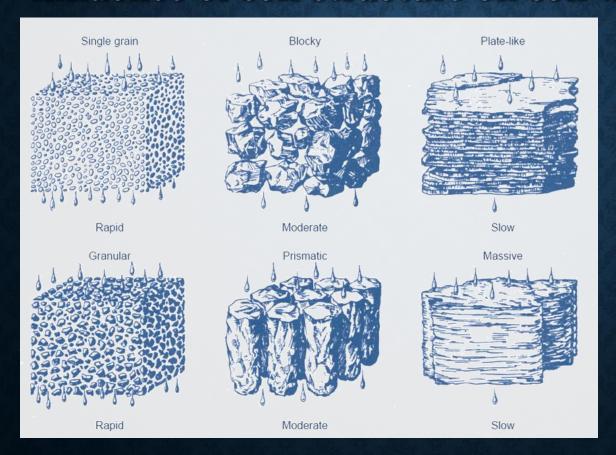






Massive

Influence of soil structure on soil water flow



Soil Density and Soil Specific Gravity

Soil density

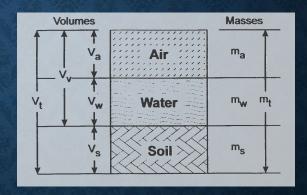
particle density

$$\rho_{s} = \frac{m_{s}}{V_{s}}$$

bulk density

$$\rho_b = \frac{m_s}{V_t}$$

$$= \frac{m_s}{V_a + V_w + V_s}$$

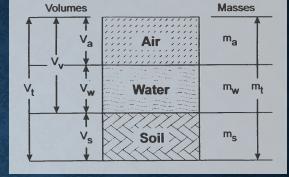


Specific gravity

real specific gravity

$$R_{S} = \frac{\rho_{S}}{\rho_{W}}$$

$$= \frac{m_{S}}{V_{S} \times \rho_{W}}$$



apparent specific gravity

$$A_{s} = \frac{\rho_{b}}{\rho_{w}}$$

$$= \frac{m_{s}}{V_{t} \times \rho_{w}}$$

Pore Space

Porosity

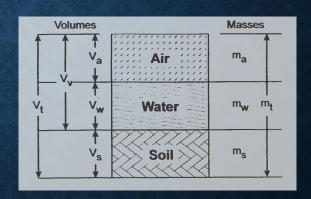
Void ratio

Pore space

• Liquid or gas components between soil particle

$$\bullet \qquad \qquad = \qquad \frac{V_{v}}{V_{t}}$$

$$e = \frac{V_{v}}{V_{s}}$$



Porosity of soil

เนื้อคิน	ความพรุน (%)		
	ช่วงค่าปกติ	ค่าเฉลี่ย	
Sand	32 – 42	38	
Sandy Loam	40 – 47	43	
Loam	43 – 49	46	
Clay Loam	47 – 51	49	
Silty Clay	49 – 53	51	
Clay	51 - 55	53	

Soil Water

Soil properties

Water retention in soil

Infiltration

Water Retention in Soil

Soil Water Content

- Gravimetric water content
- Volumetric water content
- Degree of saturation

Soil Water Content

Soil moisture by mass

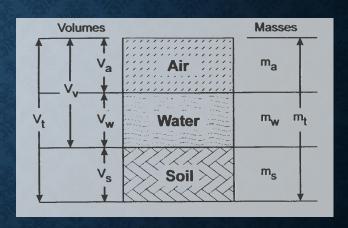
$$\theta_m = \frac{m_w}{m_s}$$

Soil moisture by volume

$$\theta_{v} = \frac{V_{w}}{V_{t}}$$

Degree of saturation

$$s = \frac{V_w}{V_v}$$



Soil core sampler





$$\rho_w \theta_v = \rho_b \theta_m
\theta_v = \frac{\rho_b}{\rho_w} \theta_m
\theta_v = A_s \theta_m$$

Water content by mass vs by volume

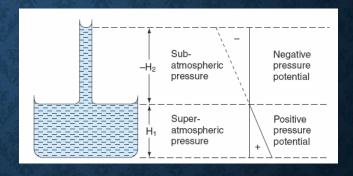
$$d = \theta_{v}D$$
$$= \theta_{m}A_{s}D$$

Depth of water in soil column

Potential Energy of Water in Soil

Soil water potential

- Gravitational potential
- Osmotic potential
- Matric potential
- Hydrostatic pressure potential

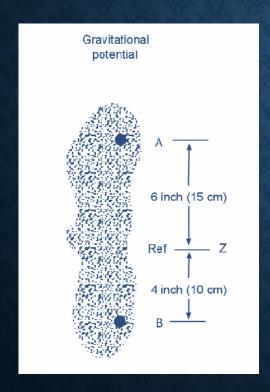


Unit

หน่วย	สัญลักษณ์	ชื่อ	มิดิ	หน่วย	หน่วย
				(CGS)	(SI)
พลังงาน/มวล	μ_T	Chemical potential	L^2/T^2	g^{-1}	$\rm Jkg^{-1}$
พลังงาน/ปริมาตร	ψ_T	Soil water potential	M/LT^2	cm^{-3}	$J m^{-3} = N m^{-2} = Pa$
พลังงาน/น้ำหนัก	h_T	Soil water potential head	L	cm	$J kg^{-1} m^{-1} s^{-2} = m$

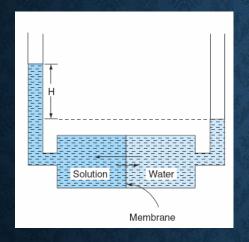
- งานทางด้านชลประทาน/แหล่งน้ำ นิยมใช้
 - พลังงานต่อน้ำหนัก (น้ำ) => เฮด (head)

Gravitational potential

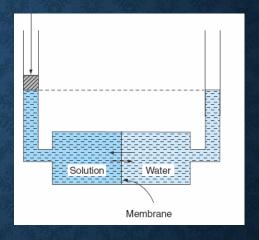


Gravitational potential

Osmotic potential

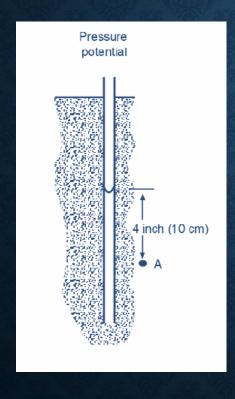


Osmosis



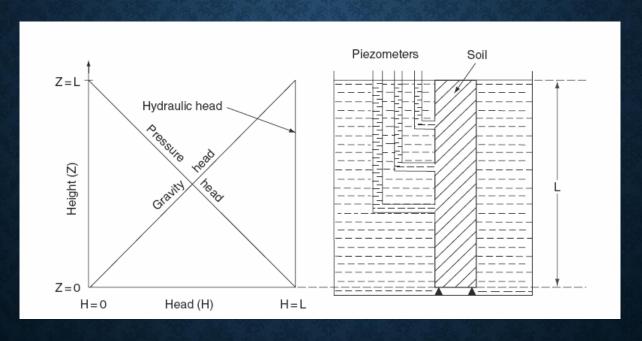
Osmotic pressure

Pressure potential



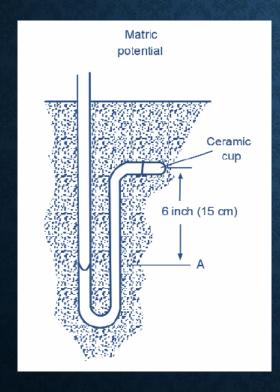
- Pressure potential
- In case of saturated soil

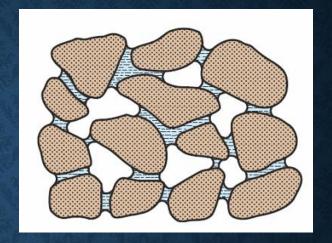
Pressure potential in saturated soil



- Pressure head
- Gravitational head

Unsaturated soil





Matric potential

Soil water potential

- soil water potential
 - **■** Energy/volume

$$\psi_T = \psi_z + \psi_s + \psi_m + \psi_p + K$$

■ head

$$h_T = z + s + h + p + K$$

Energy/weight (of water)

Saturated soil

$$h_T = z + s + p$$

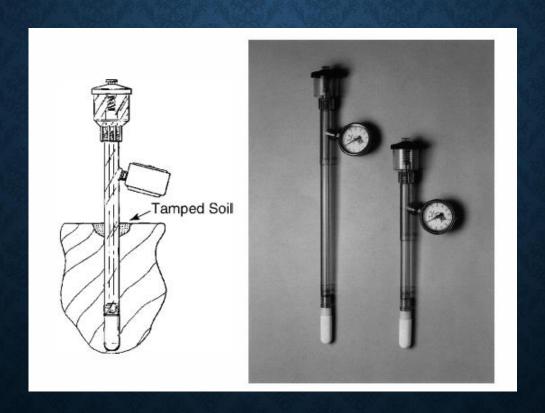
Unsaturated soil

$$h_T = z + s + h$$

Soil Moisture Tension

- Sometimes called "suction"
- Unit of measurement
 - Pascal (Pa) or N/m2
 - bar
 - Head or depth of water in m (or cm)

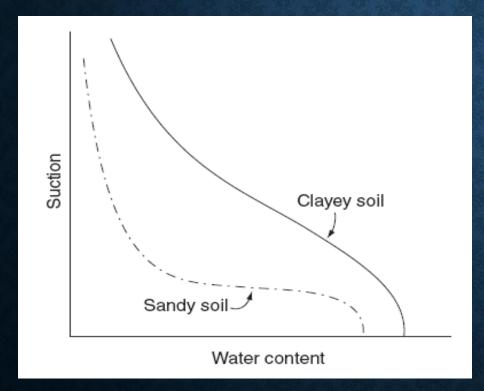
Tensiometer



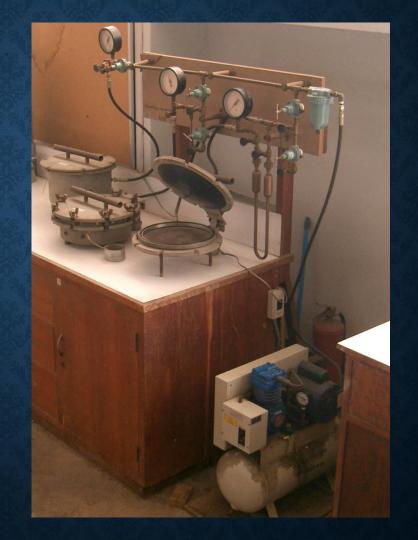


Soil-Moisture Characteristic Curve

Soil-Moisture Characteristic Curve



• Relation between (volumetric) water content and matrix potential (h)



Stages of Water Content in Soil

Water content at field capacity (FC)

Water content at permanent wilting point (PWP)

Available water

- TAW: total available water
- RAW: readily available water

Stages of Water Content in Soil

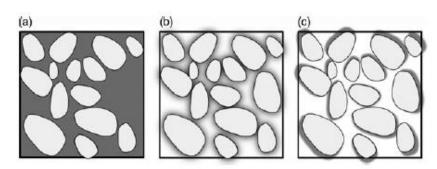


Fig. 5.4. Soil moisture at different stages of moisture content. (a) Saturation – all pores filled with water, little or no air; this situation occurs during and immediately following irrigation or rainfall. (b) Field capacity – water is held in the soil after surplus has drained away under gravitational forces. (c) Permanent wilting point – water attached by surface tension forces to soil particles, cannot be removed by plant root suction.

- From an irrigation point of view, there are different levels of water content in the soil, and four terms are used to identify these water content levels:
 - saturation
 - field capacity
 - permanent wilting point
 - · available soil water

Saturation



- During and immediately after irrigation all the pore space in the soil is filled with water and the soil is saturated.
- There is little air in the soil, and for most crops (other than rice) if the soil stays saturated the crop will be damaged due to this lack of air for the roots to breathe.
- If there are no drainage problems, the water in the soil will drain away under gravity following irrigation, leaving space for air in the soil's pore space

Field Capacity, FC

- Field capacity is the quantity of water held in the soil once the water has drained away from the saturated soil.
- This water is held to the soil particles by surface tension forces, and much of it is available for taking up by the plant's roots.
- The volume of water held by the soil at field capacity depends primarily on its texture and structure.
- The forces holding the water in the soil against the gravitational pull are surface tension forces. Soils with small particle size, such as silts and clays, have a large surface area and thus can hold more water.

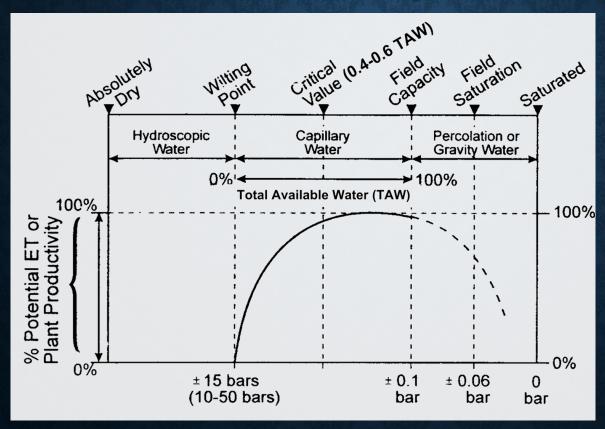
Permanent Wilting Point, WP

- by the plant's roots exerting a greater pull or tension than the surface tension holding the water to the soil particle's surface.
- At some point, termed the permanent wilting point,
 the suction exerted by the plant's roots is not sufficient to remove the water from around the soil's particles.
- At this point, the crop will become stressed, the yield will be reduced and the crop
 may perish. At permanent wilting point a crop's leaves may droop or wilt; in some
 crops, there will be a change in appearance in the leaf color.
- The permanent wilting point is affected by the soil texture in the same way as with field capacity, thus for fine-textured soils the moisture content at permanent wilting point is higher than for coarse-textured soils.

Available Soil Moisture (FC - WP)

- The water available to the plant is the difference between the moisture content at field capacity and that at the permanent wilting point.
- Though there may still be water in the soil at the permanent wilting point it cannot be removed by the plant, and is thus unavailable.
- The objective of irrigation is to allow the soil moisture to reduce to a safe limit
 (above the permanent wilting point) and then to irrigate the soil to bring it back to
 field capacity.
- The interval between irrigation will thus depend on the available moisture in the soil and the rate at which the soil water is abstracted by the crop

Soil water content vs ET/Productivity



Soil Water

Soil properties

Water retention in soil

Infiltration

Infiltration

Definition
Infiltration equations
Infiltration in Irrigation

• saturation zone

- เริ่มจากส่วนที่อยู่ใกล้กับผิวดิน
- จะมีความลึกเพิ่มขึ้นเรื่อย ๆ

• transmission zone

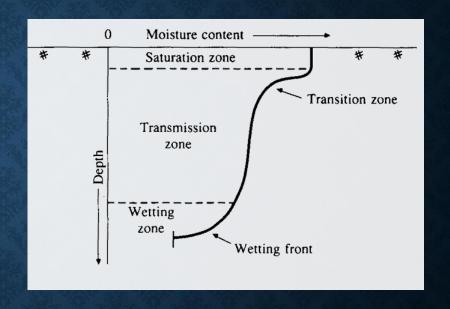
- เป็นส่วนที่น้ำไหลผ่านชั้นดินที่ยังไม่อิ่มตัว
- ปริมาณความชื้นตลอดหน้าตัดใกล้เคียงกัน

wetting zone

- เป็นส่วนที่ความชื้นกำลังเพิ่มขึ้นอย่างรวดเร็ว
- ในชั้นดินที่ลึกลงไป ในช่วงเริ่มต้นยังมีความชื้นน้อย

wetting front

- เป็นหน้าตัดที่เริ่มเปียกน้ำและกำลังมีการเปลี่ยนแปลงความชื้นอย่างรวดเร็ว
- ดินมีความซึ้นแตกต่างกันมากจนแยกระหว่างดินเปียกกับดินแห้งได้อย่างชัดเจน

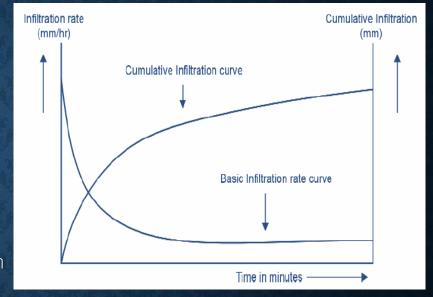


• Infiltration rate (f)

- อัตราการซึมมีค่าสูงในช่วงแรก
- ค่อย ๆ ลดลงเมื่อเวลาผ่านไป
 จนกระทั่งคงที่

Cumulative infiltration (F)

- ในขณะเริ่มต้น ปริมาณน้ำที่ซึมผ่านผิวดินจะเป็นศูนย์
- มีค่าเพิ่มขึ้นตามเวลา
- โดยในช่วงแรก โค้งปริมาณการซึมผ่านผิวดินมีความชั้นมาก
- เมื่อเวลาผ่านไป ความชันของโค้งจะลดลง
- ในที่สุด กราฟปริมาณการซึมผ่านผิวดินจะเป็นเส้นตรง เนื่องจาก อัตราการซึมผ่านผิวดินคงที่



Relation between infiltration rate and cumulative infiltration

$$F(t) = \int_0^t f(t)dt$$

$$f(t) = \frac{dF(t)}{dt}$$

- อินทิเกรต สมการอัตราการซึมผ่านผิวดิน (f) ได้ สมการปริมาณการซึมผ่านผิวดิน (F)
- หาอนุพันธ์ สมการปริมาณการซึมผ่านผิวดิน **(F)** ได้ สมการอัตราการซึมผ่านผิวดิน (**f**)

Infiltration equation

$$F(t) = \gamma t^a$$

$$f(t) = a\gamma t^{a-1}$$

 Lewis's equation or Kostiakov's equation

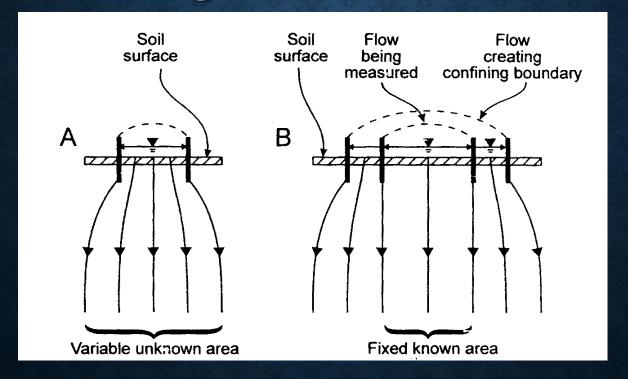
Double-ring infiltrometer





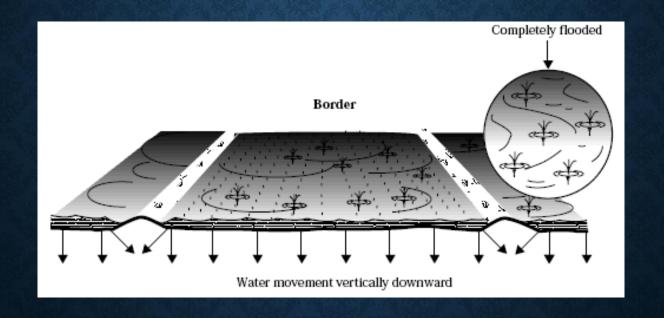


Infiltration measurement by double-ring infiltrometer

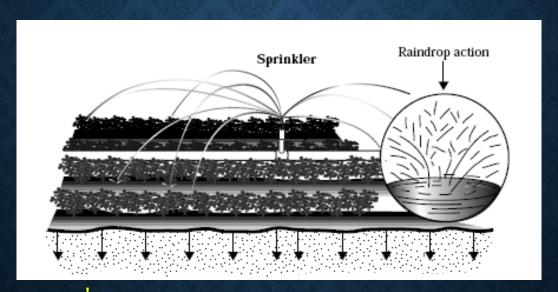


Infiltration in Irrigation

Infiltration in border irrigation

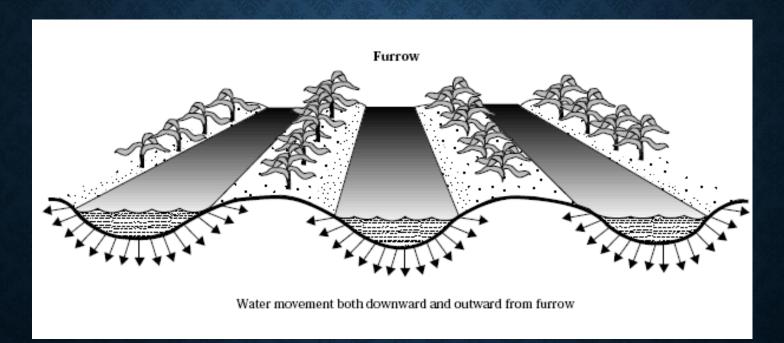


Infiltration in sprinkler irrigation

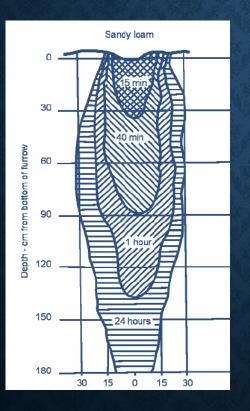


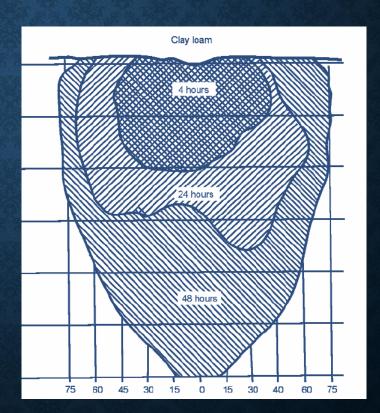
- น้ำซึมทางดิ่งอย่างเดียว
- ■อัตราการให้น้ำต้องไม่เกินอัตราการซึมผ่านผิวดิน

Infiltration in furrow irrigation

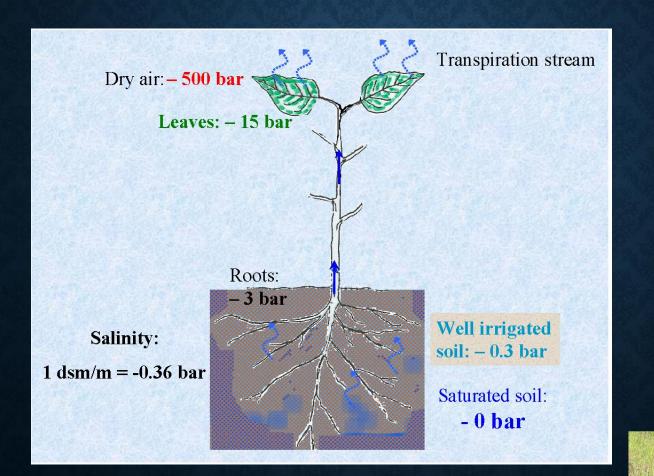






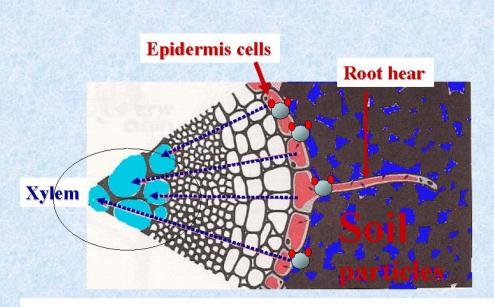


Water Flow in Plant



Shlomo Kramer

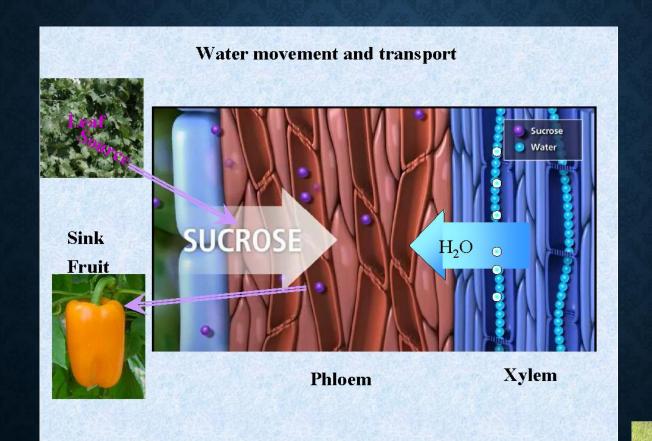
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Water passing by <u>OSMOSIS</u> from cell to cell till arriving to the xylem tubes

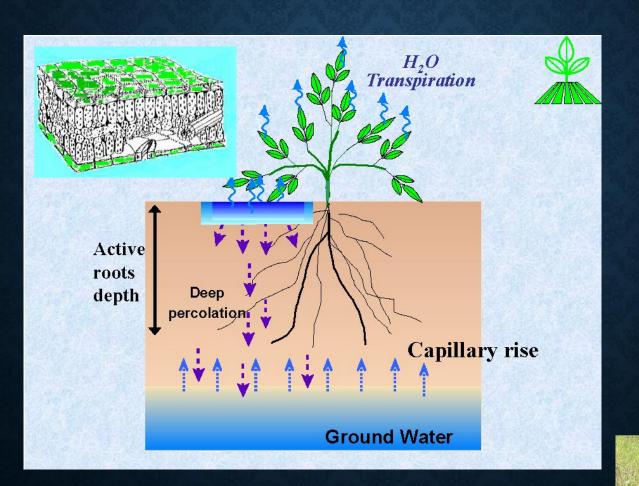
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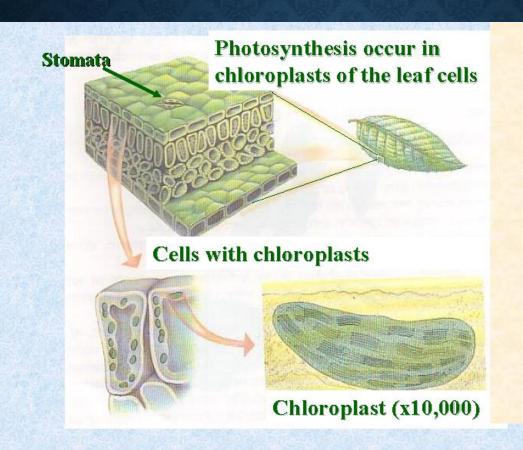
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Gas exchange through the Stomata opening



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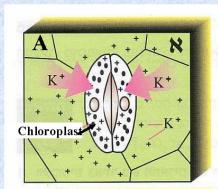
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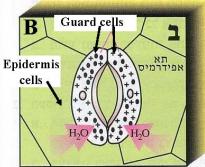


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Stomata opening - closing





 ${\rm CO_2}$ required for photosynthesis diffuse into the leaf through the open stomata.

Water molecules diffuse out in transpiration process.

<u>Factors affect the opening – closing process:</u>

- · Light intensity
- · Water status in the plant.
- · CO, concentration in the leaf
- ·Opening at light time,
- •* Closing at dark, or when the plant

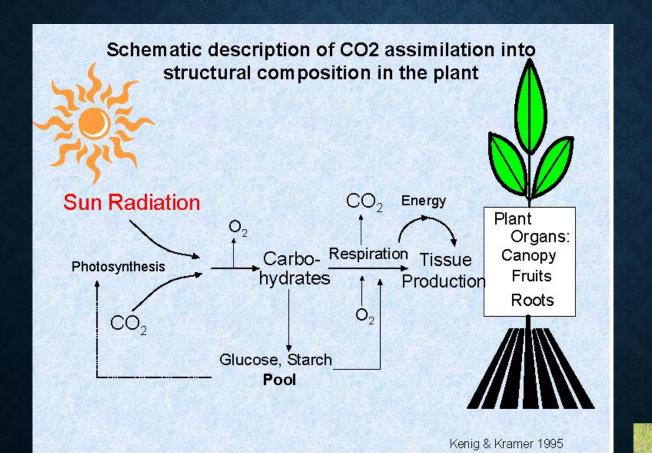
is in stress situation

Light + low $\rm CO_2 > K^+$ entering into guard cells > osmotic pressure increased > water osmosis to guard cell, turgor become higher> stomata open

Dark + high CO₂ > K⁺> stomata close

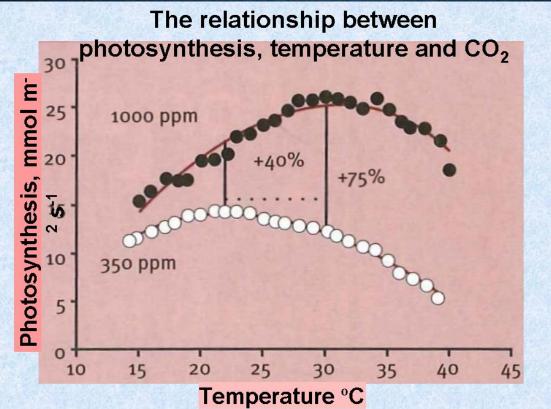
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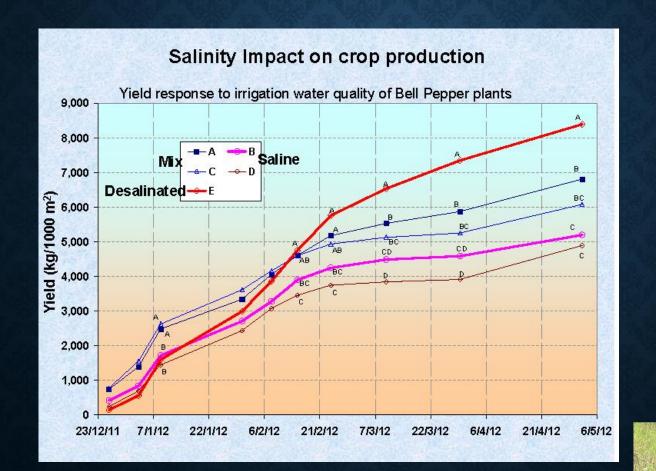
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After: Peter van Weel, PPO greenhouse horticulture, Wageningen, the Netherlands

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Thank you